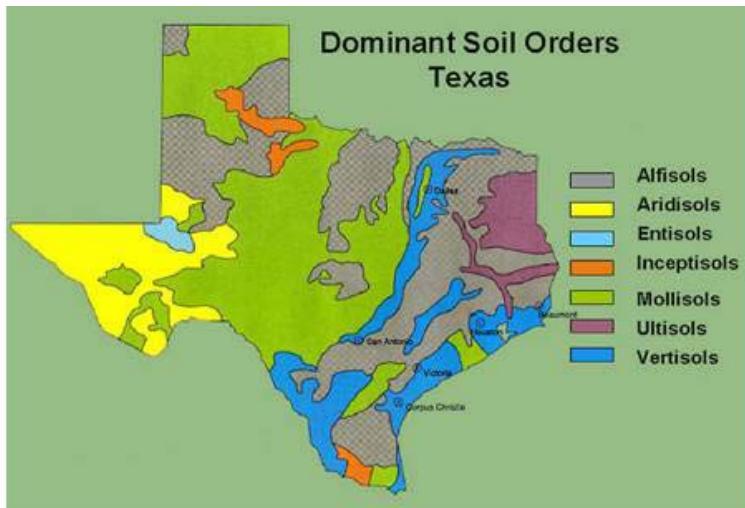
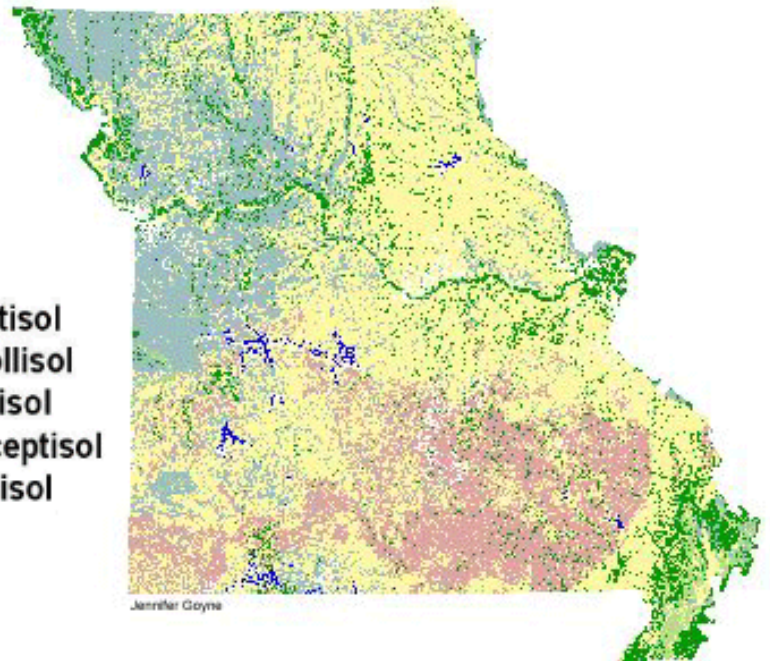


Name _____

- 1)
 - a. Using the soil taxonomy poster, surface map, and geologic maps, what can you hypothesize about the classification and diversity of Indiana soils?
 - b. What about the soils of your home state/region? (If your home state is IN, hypothesize about another location!)
- 2) Explain the soil order distribution of
 - a. Texas
 - b. Missouri



getting a feel for soils



- 3) Use the guide and flow chart to determine the textural class and approximate sand, silt, and clay proportions of the following samples:
 - a. unknown light material
 - b. Planting Experiment Soil 1
 - c. Planting Experiment Soil 2
 - d. Planting Experiment Soil 3
- 4) Do you think you might get different answers if you performed the analysis again? Why or why not?
- 5) According to your own experience, and the definitions discussed in class and in your text, are these materials *soils*? Explain your answer.
- 6) Do you think testing the same soil by feel and by quantitative/mechanical analysis (e.g., sieve, hydrometer, laser diffraction, etc.) might yield different textural classes for any of these samples? Why or why not? If so, which assessment would you consider more v

Bedrock Geology

PENNSYLVANIAN

- McLeansboro Group
shale, sandstone, limestone,
and thin coals
- Carbondale Group
shale, sandstone, limestone,
and thick coals
- Raccoon Creek Group
sandstone, shale, clay,
limestone, and thin coals

MISSISSIPPIAN

- Buffalo Wallow, Stephensport, and West Baden Groups
shale, sandstone, and limestone
- Blue River and Sanders Groups
limestone and dolomite
- Borden Group and Rockford Limestone
shale, siltstone, and limestone
- Coldwater Shale
gray shale

DEVONIAN

- Ellsworth and Sunbury Shales
gray, green, and black shale
- New Albany Shale
black shale
- Antrim Shale
black shale
- Muscatatuck Group
limestone and dolomite

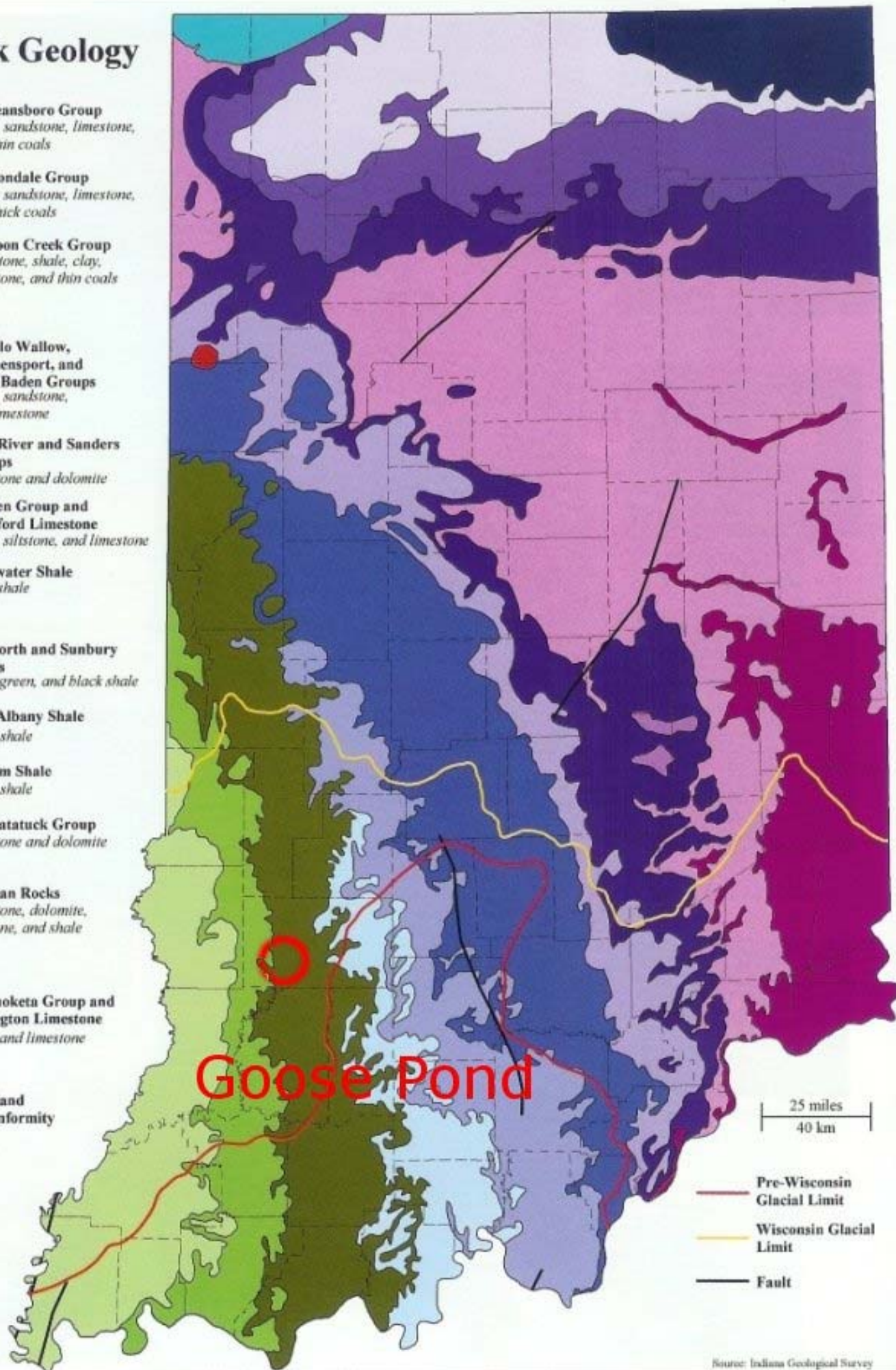
SILURIAN

- Silurian Rocks
limestone, dolomite, siltstone, and shale

ORDOVICIAN

- Maquoketa Group and Lexington Limestone
shale and limestone

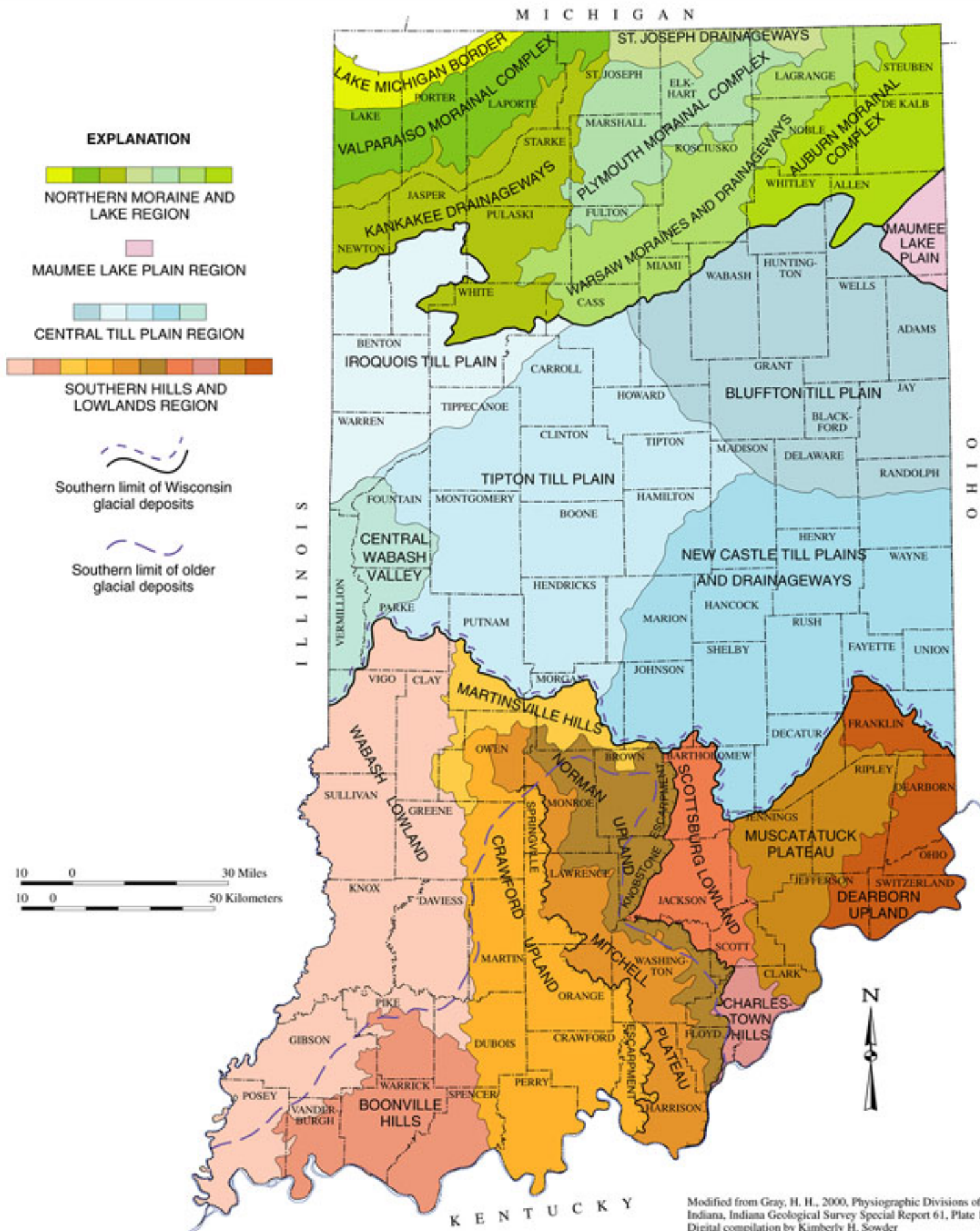
- Kentland Unconformity



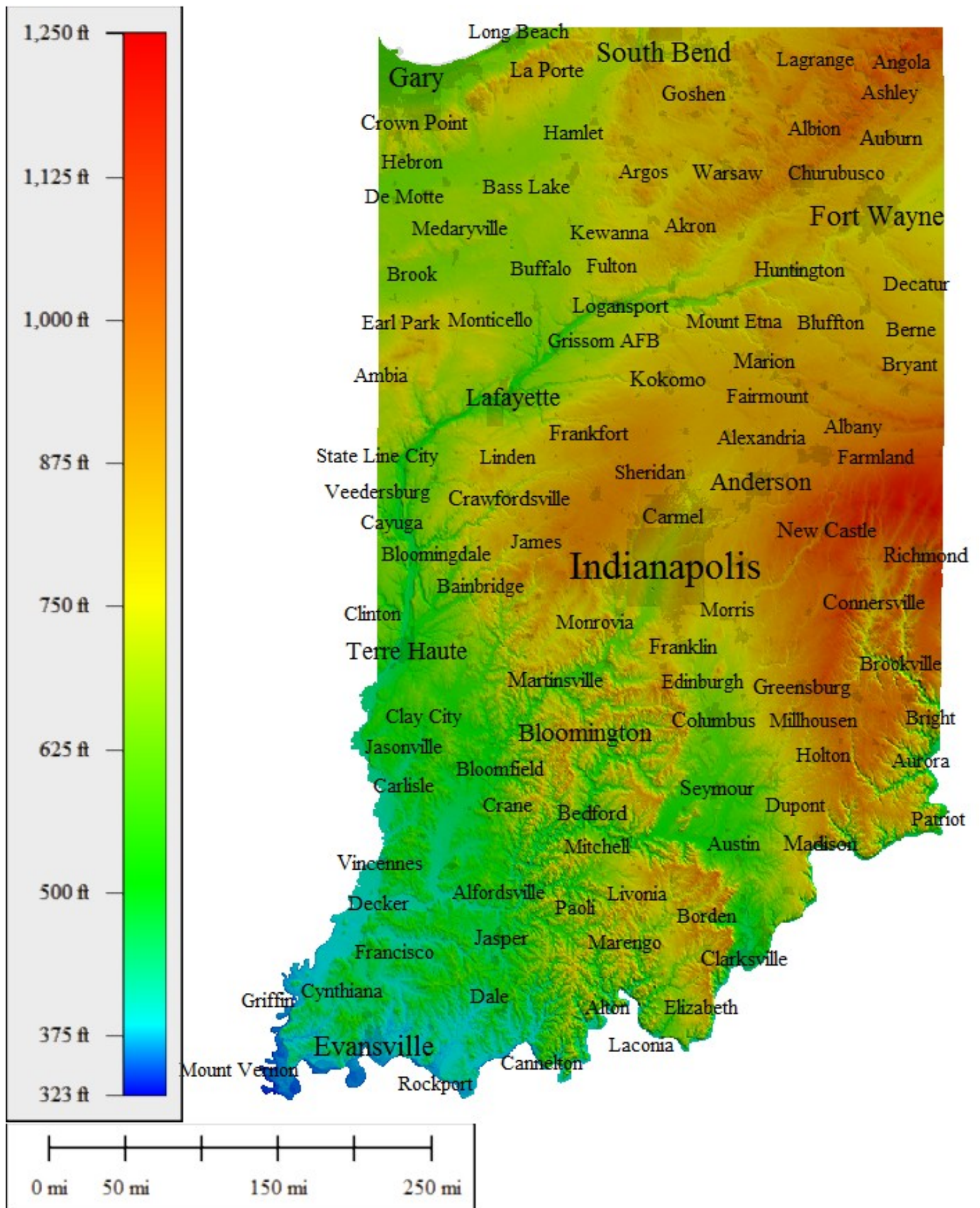
Goose Pond

25 miles
40 km

- Pre-Wisconsin Glacial Limit
- Wisconsin Glacial Limit
- Fault



Modified from Gray, H. H., 2000, Physiographic Divisions of Indiana, Indiana Geological Survey Special Report 61, Plate 1. Digital compilation by Kimberly H. Sowder





ALFISOLS

Alfisols are found in moist areas. These soils result from weathering processes that leach clay minerals and other constituents out of the surface layer and into the subsoil, where they can build and supply moisture and nutrients to plants. They formed primarily under forest or mixed vegetative cover and are productive for most crops.

ALFISOLS MAKE UP ABOUT 10% OF THE WORLD'S ICE-FREE LAND SURFACE.



ANDISOLS

Andisols have been weathering processes that generate minerals with little orderly crystalline structure. These minerals can result in an unusually high water- and nutrient-holding capacity. As a group, Andisols tend to be highly productive soils. They include weakly weathered soils with much volcanic glass as well as more strongly weathered soils. They are common in cool areas with moderate to high precipitation, especially those areas associated with volcanic materials.

ANDISOLS MAKE UP ABOUT 1% OF THE WORLD'S ICE-FREE LAND SURFACE.



ARIDISOLS

Aridisols are soils that are too dry for the growth of mesophytic plants. The lack of moisture greatly restricts the intensity of weathering processes and limits most soil development processes to the upper part of the soil. Aridisols often contain gypsum, salt, calcium carbonate, and other materials that are easily leached from soils in more humid environments.

ARIDISOLS MAKE UP ABOUT 12% OF THE WORLD'S ICE-FREE LAND SURFACE.



ENTISOLS

Entisols are soils that show little or no evidence of pedogenic horizon development. Entisols occur in areas of recently deposited parent materials or in areas where erosion or deposition rates are faster than the rate of soil development, such as dunes, steep slopes, and flood plains. They occur in many environments.

ENTISOLS MAKE UP ABOUT 16% OF THE WORLD'S ICE-FREE LAND SURFACE.



GELISOLS

Gelisols are soils that have formed near the soil surface under low evidence of cryoturbation (frost churning) and/or ice segregation. Gelisols are common in the higher latitudes or at high elevations.

GELISOLS MAKE UP ABOUT 9% OF THE WORLD'S ICE-FREE LAND SURFACE.



HISTOSOLS

Histosols have a high content of organic matter and no gramine. They are saturated year-round, but a few are freely drained. Histosols are commonly called bogs, meadows, peats, or mucks.

Histosols have in decomposed plant remains that accumulate in water, forest floor, or more under than they decay. If these soils are drained and exposed to air, microbial decomposition is accelerated and the soils may subside dramatically.

HISTOSOLS MAKE UP ABOUT 1% OF THE WORLD'S ICE-FREE LAND SURFACE.



MOLLISOLS

Mollisols are soils that have a dark colored surface horizon relatively high in content of organic matter. The soils are bare with throughout and therefore are quite fertile.

Mollisols characteristically form under grass to climates that have a moderate to pronounced seasonal moisture deficit. They are extensive soils on the steppes of Europe, Asia, North America, and South America.

MOLLISOLS MAKE UP ABOUT 7% OF THE WORLD'S ICE-FREE LAND SURFACE.



OXISOLS

Oxisols are highly weathered soils of tropical and subtropical regions. They are dominated by low activity minerals, such as quartz, kaolinite, and iron oxides. They tend to have indistinct horizons.

Oxisols characteristically occur on land surfaces that have been stable for a long time. They have low natural fertility as well as a low capacity to retain additions of lime and fertilizer.

OXISOLS MAKE UP ABOUT 8% OF THE WORLD'S ICE-FREE LAND SURFACE.



SPODOSOLS

Spodosols formed from weathering processes that strip organic matter combined with aluminum leach or without iron leach the surface layer and deposit them in the subsoil. In undisturbed areas, a gray eluvial horizon that has the color of weathered quartz, kaolinite, and iron oxides.

Spodosols commonly occur in areas of coarse-textured deposits under coniferous forests of humid regions. They tend to be acid and infertile.

SPODOSOLS MAKE UP ABOUT 4% OF THE WORLD'S ICE-FREE LAND SURFACE.



ULTISOLS

Ultisols are soils in humid areas. They formed from fairly intense weathering and leaching processes that result in a clay-enriched subsoil dominated by minerals, such as quartz, kaolinite, and iron oxides.

Ultisols are typically acid soils in which most nutrients are concentrated in the upper few inches. They have a moderately low capacity to retain additions of lime and fertilizer.

ULTISOLS MAKE UP ABOUT 8% OF THE WORLD'S ICE-FREE LAND SURFACE.



INCEPTISOLS

Inceptisols are soils of semiarid to humid environments that generally exhibit only moderate degrees of soil weathering and development.

Inceptisols have a wide range in characteristics and occur in a wide variety of climates.

INCEPTISOLS MAKE UP ABOUT 17% OF THE WORLD'S ICE-FREE LAND SURFACE.



VERTISOLS

Vertisols have a high content of expanding clay minerals. They undergo pronounced changes in volume with changes in moisture. They have cracks that open and close periodically and that show evidence of soil movement in the profile.

Because they swell when wet, vertisols transmit water very slowly and have underground little leaching. They tend to be fairly high in natural fertility.

VERTISOLS MAKE UP ABOUT 2% OF THE WORLD'S ICE-FREE LAND SURFACE.



Entisols are soils that show little or no evidence of pedogenic horizon development. Entisols occur in areas of recently deposited parent materials or in areas where erosion or deposition rates are faster than the rate of soil development, such as dunes, steep slopes, and flood plains. They occur in many environments.

Entisols are common in the deserts of the world.

ENTISOLS MAKE UP ABOUT 16% OF THE WORLD'S ICE-FREE LAND SURFACE.



Gelisols are soils that have formed near the soil surface under low evidence of cryoturbation (frost churning) and/or ice segregation. Gelisols are common in the higher latitudes or at high elevations.

Gelisols are common in the deserts of the world.

GELISOLS MAKE UP ABOUT 9% OF THE WORLD'S ICE-FREE LAND SURFACE.



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HISTOSOLS MAKE UP ABOUT 1% OF THE WORLD'S ICE-FREE LAND SURFACE.

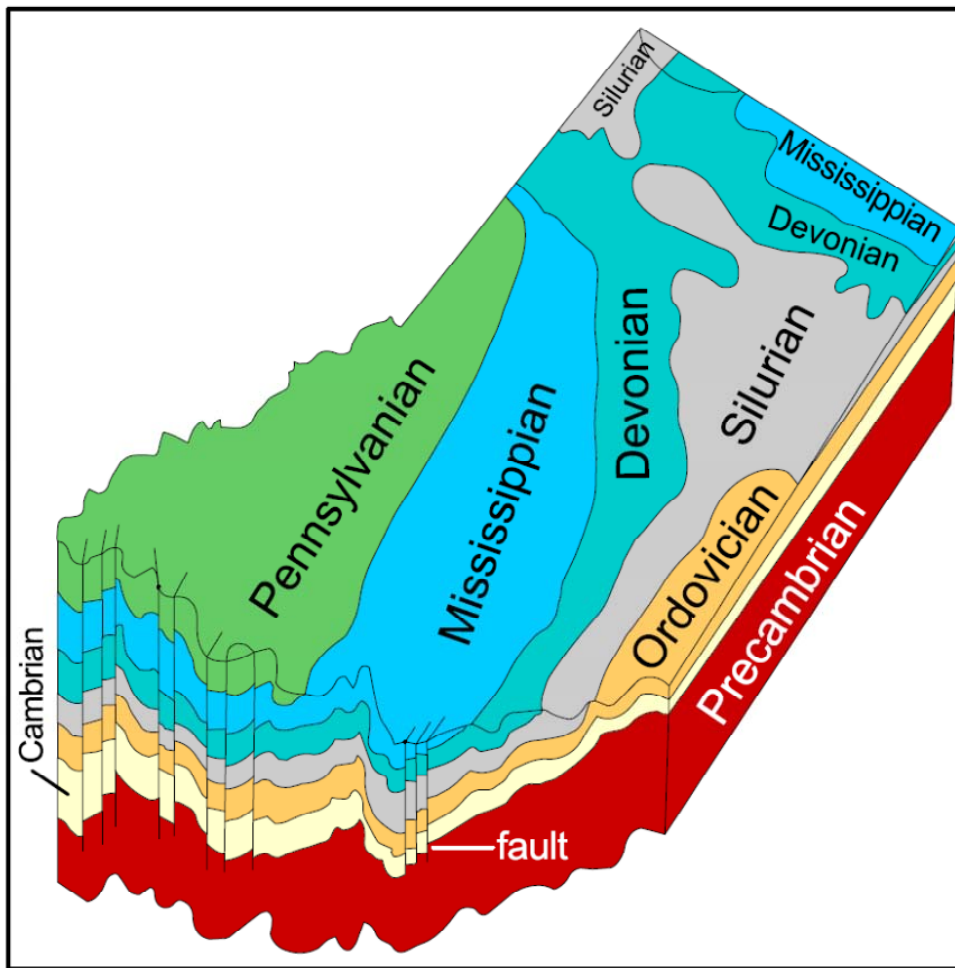


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MOLLISOLS MAKE UP ABOUT 7% OF THE WORLD'S ICE-FREE LAND SURFACE.

THE TWELVE ORDERS OF SOIL TAXONOMY



Block diagram illustrating southwesterly dip of the Sedimentary rocks that make up the bedrock in Indiana. Notice the outcrop and subcrop pattern that has resulted from erosion on the gently dipping bedrock surface.

Guide to Texture by Feel

Modified from:
Thien, S.J. (1979). A flow diagram for teaching texture by feel analysis.
Journal of Agronomic Education.
8:54-55.

Texture class is one of the first things determined when a soil is examined. It is related to weathering and parent material. The differences in horizons may be due to the differences in texture of their respective parent materials.

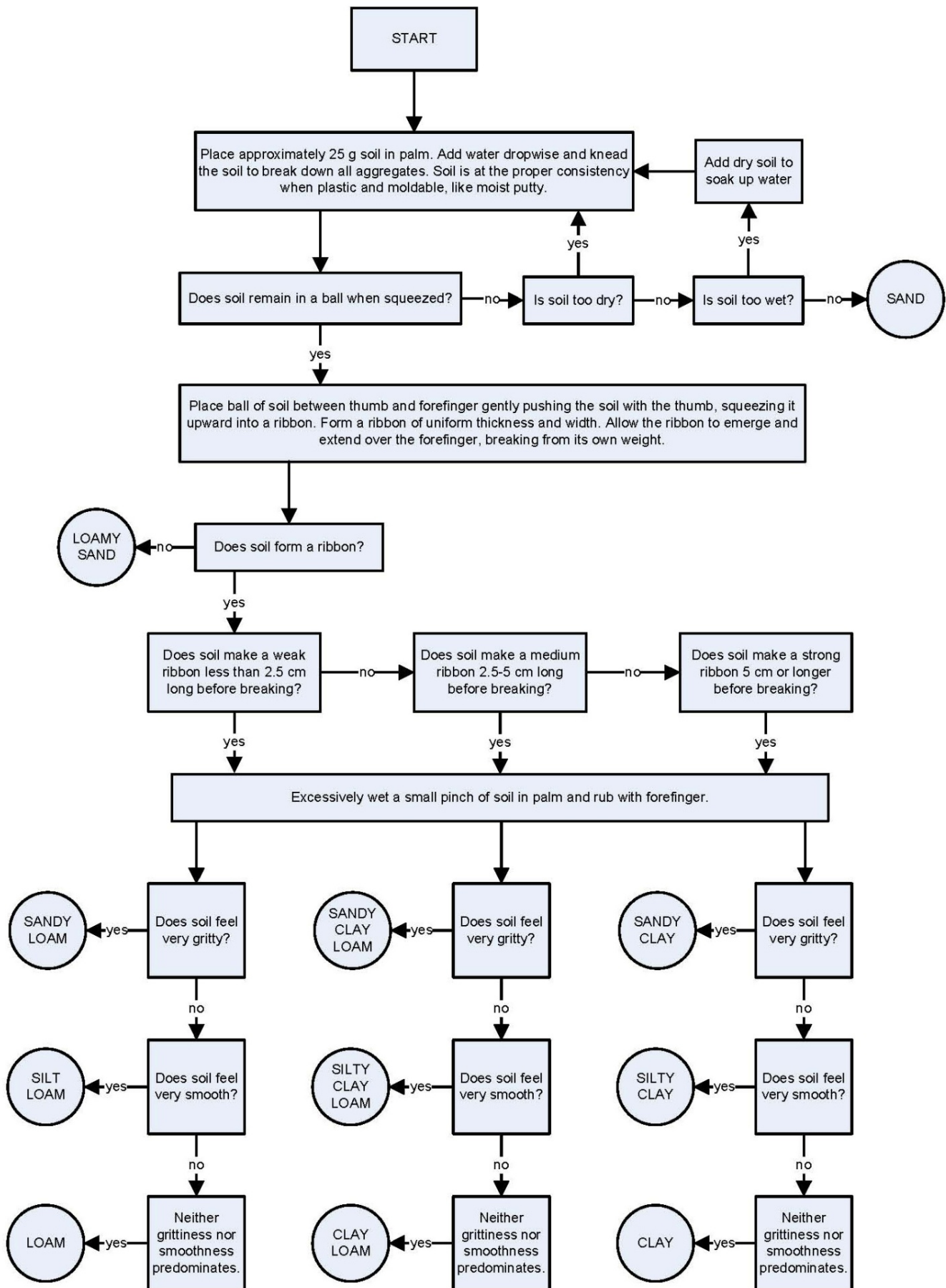
Texture class can be determined fairly well in the field by feeling the sand particles and estimating silt and clay content by flexibility and stickiness. There is no field mechanical-analysis procedure that is as accurate

as the fingers of an experienced scientist, especially if standard samples are available. A person must be familiar with the composition of the local soils. This is because certain characteristics of soils can create incorrect results if the person does not take these characteristics into account.

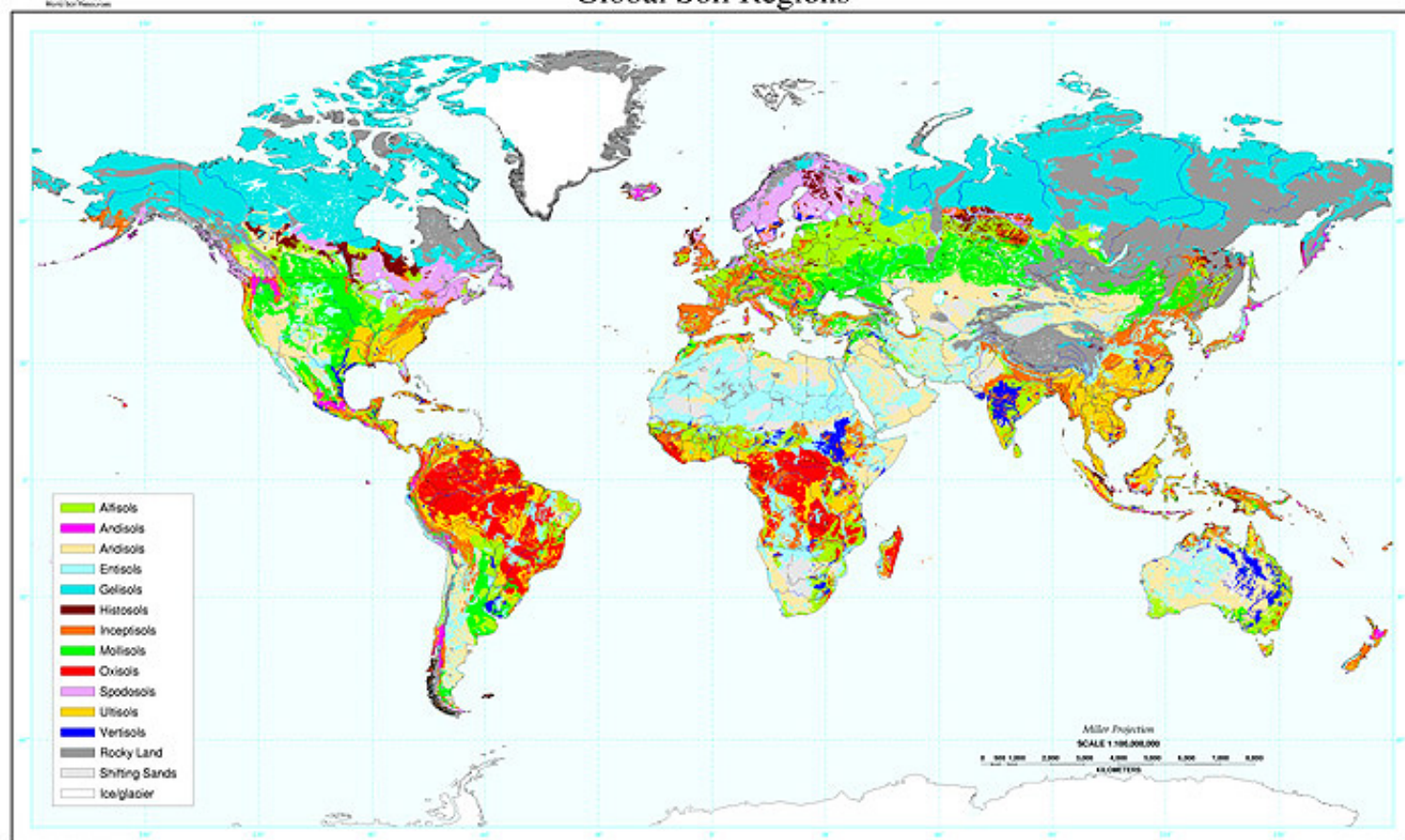
In some environments clay aggregates form that are so strongly cemented together that they feel like fine sand or silt. In humid climates iron oxide is the cement. In desert climates silica is the cement and in arid regions lime can be the cement. It takes prolonged rubbing to show that they are clays and not silt loams.

Some soils derived from granite contain grains that resemble mica but are softer. Rubbing breaks down these grains and reveals that they are clay. These grains resist dispersion and field and laboratory determinations may disagree.

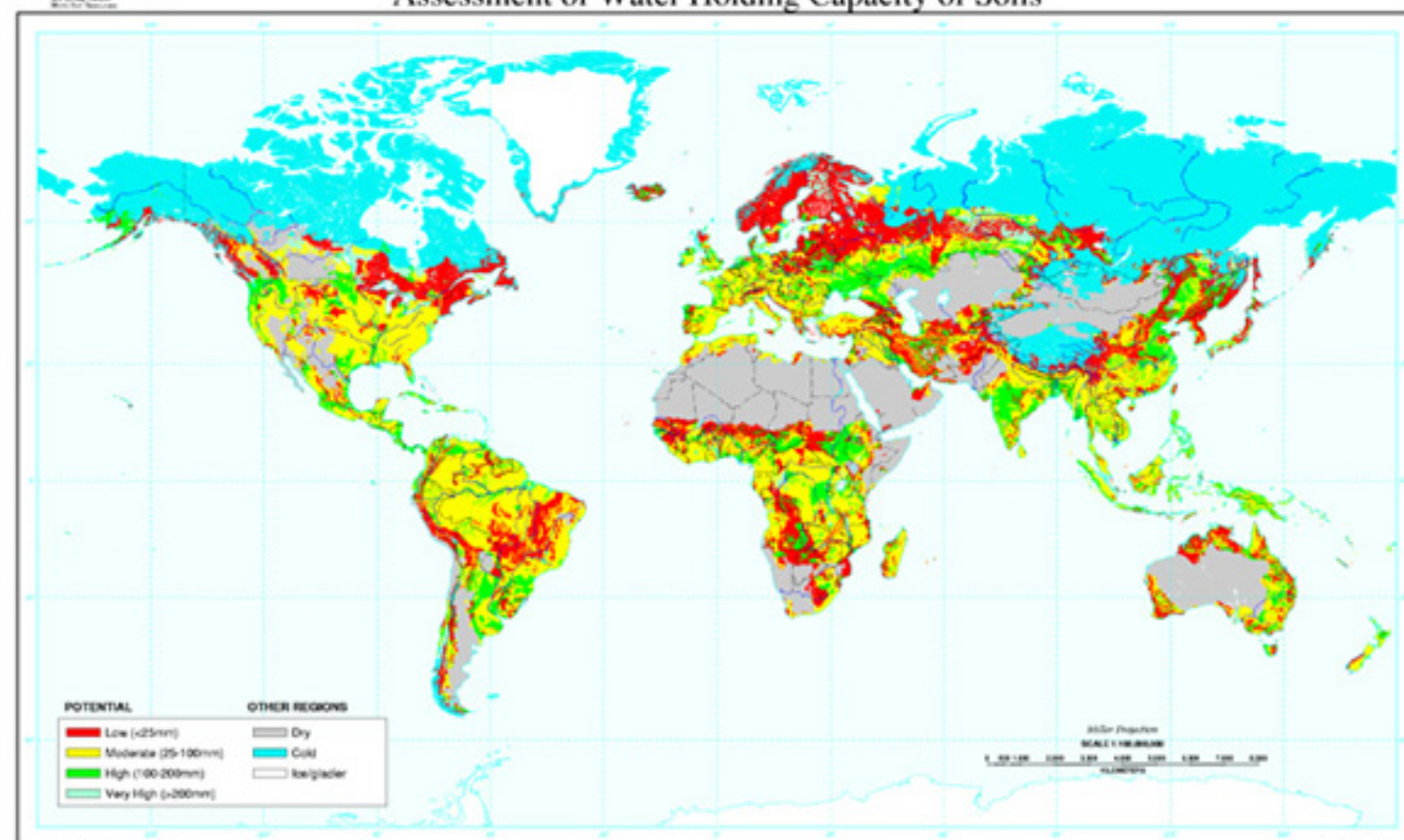
Many soil conditions and components mentioned earlier cause inconsistencies between field texture estimates and standard laboratory data. These are, but not limited to, the presence of cements, large clay crystals, and mineral grains. If field and laboratory determinations are inconsistent, one or more of these conditions is suspected.

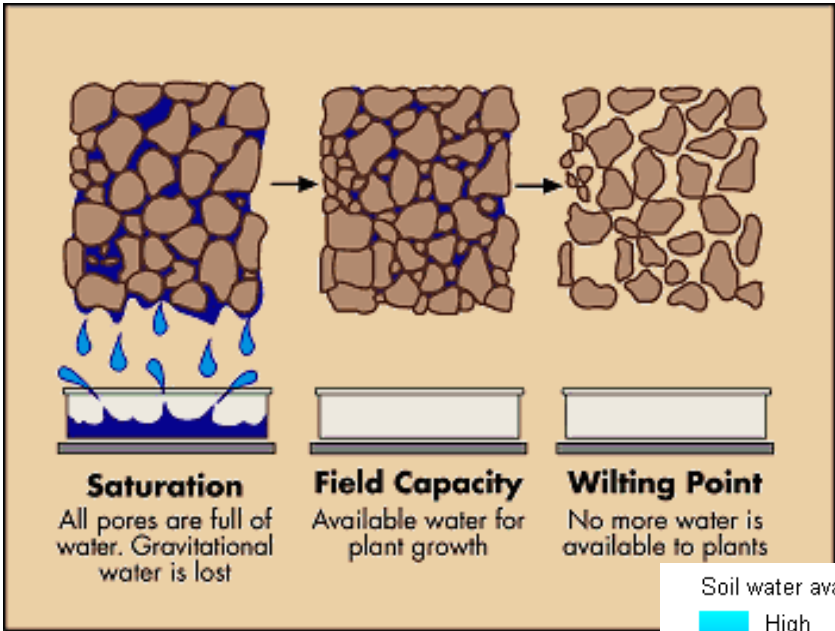


Global Soil Regions

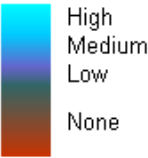


Assessment of Water Holding Capacity of Soils





Soil water availability to plant



Soil Textural Triangle

